

Article

Research on Inversion Log Evaluation Method of Special Mineral in Alkali Lake Shale Oil Reservoir—A Case Study of the Fengcheng Formation in the Mahu Sag, China

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Abstract: The Fengcheng Formation in the Mahu Sag of the Junggar Basin, China, is characterized by alkaline lake deposits, featuring abundant alkaline minerals. The content of alkaline minerals affects the physical properties and oil-bearing properties of the reservoir, and existing mineral inversion methods cannot calculate the content of alkaline minerals. Based on Litho Scanner Log data, we can calculate the dry weight of elements using the oxide closure model. By improving the rock volume physical model; adding trona, shortite, eitelite, and reedmergnerite to the rock volume physical model; and combining with the least squares method, the mineral content calculation was carried out, using the inversion method of combination models (Shortite Model, Eitelite Model, Reedmergnerite Model, and Trona Model) to achieve mineral inversion of alkali-bearing shale oil reservoirs. Litho Scanner Log is expensive, and its widespread application will increase exploration costs. This article scales the mineral inversion results of Litho Scanner Log into conventional log data, improves the rock volume physical model of conventional log, and uses a combination model to achieve mineral inversion of alkali-bearing shale oil reservoirs in conventional log. Compared with the results of X-ray diffraction experiments, the average absolute error of all minerals except for trona and feldspar is less than 10%, and the inversion results are consistent with the core test results. The research results of this article can provide theoretical and technical support for the log evaluation of alkali-bearing shale oil reservoirs.

Keywords: Junggar Basin; Mahu Sag; Fengcheng Formation; alkaline minerals; Litho Scanner Log; combining model; mineral inversion



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1. Introduction

In the Junggar Basin of China, the Fengcheng Formation in the Mahu Sag has undergone comprehensive exploration deployment within the framework of the entire hydrocarbon system accumulation model. The initial well testing results have been promising, showcasing significant exploration potential. The Fengcheng Formation in the Mahu Sag is an alkaline lake facies deposit [1], with complex lithology and abundant development of alkaline minerals. Alkaline minerals are easily soluble, and during the process of hydrocarbon generation and acid expulsion in the source rock, alkaline minerals are prone to react with acidic fluids in the shale, forming corrosion pores, which can become a good space for storing oil and gas. During the formation process, alkaline minerals also frequently exchange fluids with fluids in the pores. The oil and gas generated by hydrocarbon source rocks can migrate with fluid exchange, promoting the filling of oil and gas [2]. Therefore, the content of alkaline minerals affects the physical and oil-bearing properties of reservoirs. However, currently, research on alkaline minerals is mainly focused on qualitative

evaluation, and there has been no significant progress in quantitative evaluation research. In this study, based on the Litho Scanner Log data, according to the characteristic elements selected from different alkaline minerals, we established models of shortite, eitelite, and reedmergnerite, and the trona content was determined by employing the ratio of virgin zone resistivity to flush zone resistivity. By combining the models, the quantitative inversion of alkaline minerals was realized.

Litho Scanner Log can accurately measure the content of elements such as silicon, aluminum, calcium, iron, and sodium in the formation. Through coefficient matrices, optimization algorithms, and Schlumberger's Elan mineral calculation methods, the formation mineral content can be quantitatively calculated [3–5]. However, the Fengcheng Formation in the Mahu Sag is a shale oil reservoir containing alkaline minerals, with the development of carbonate minerals, sulfate minerals, halides, borosilicate minerals, etc. [6]. The current mineral calculation model does not include alkaline minerals, so the content of alkaline minerals cannot be determined. This study takes the Fengcheng Formation in Mahu Sag as the research object; based on Litho Scanner Log data, it improves the existing mineral content calculation model; and then it combines multiple mineral combination model inversion methods to achieve mineral inversion of alkali-bearing shale oil reservoirs.

2. Geological Background

The Mahu Sag is situated in the northwestern part of the Junggar Basin, representing one of the six major hydrocarbon-rich sags in the region [7], as depicted in Figure 1. To its northwest lies the Kebai Fault Belt and the Wuxia Fault Belt, while the southeastern boundary neighbors the Xiayan Uplift and the Dabasong Uplift. Previous studies have extensively investigated various aspects of the Fengcheng Formation in the Mahu Sag, including its sedimentary, structural, reservoir, and hydrocarbon generation characteristics [8,9]. Scholars assert that the deposition period of the Fengcheng Formation was a crucial phase in the development of the foreland basin on the northwestern margin of the Junggar Basin [10], giving rise to a set of high-quality source rocks. This formation serves as a significant petroleum source in the Mahu Sag. During the deposition of the Fengcheng Formation, alternating wet and arid climates prevailed, and a pronounced evaporative environment led to a lowering of the lake level, salinization of the water body, and the development of alkaline minerals. Along the basin margins and slopes, deposition primarily occurred in marginal and shallow lake environments, while the central basin witnessed sedimentation in semi-deep-to-deep lake facies [11].

The Fengcheng Formation in the Mahu Sag exhibits complex lithology, divided into three segments from bottom to top: Feng-1(P_1f^1), Feng-2(P_1f^2), and Feng-3(P_1f^3). The Feng-1 segment(P_1f^1) is characterized by thin interbeds of mudstone and fine sandstone [12]. The Feng-2 segment(P_1f^2) predominantly consists of mudstone and shale, with a notable presence of abundant alkaline minerals. The Feng-3 segment(P_1f^3) is primarily composed of terrestrial clastic rocks and basalt. The reservoir pores in the Fengcheng Formation mainly include intergranular pores, intercrystalline pores, and microfractures. The three main pore structure types are the dissolution-pore–microfracture type, dissolution-pore–bedding-fracture type, and mesopore–bedding-fracture type. The pores in the Fengcheng Formation are small, with a diameter mainly distributed between 30 and 200 nm [13], exhibiting poor pore connectivity. The reservoir porosity ranges from 0.2% to 13.7%, and permeability varies from 0.01 mD to 6.85 mD, categorizing it as a tight, low-permeability reservoir [14].

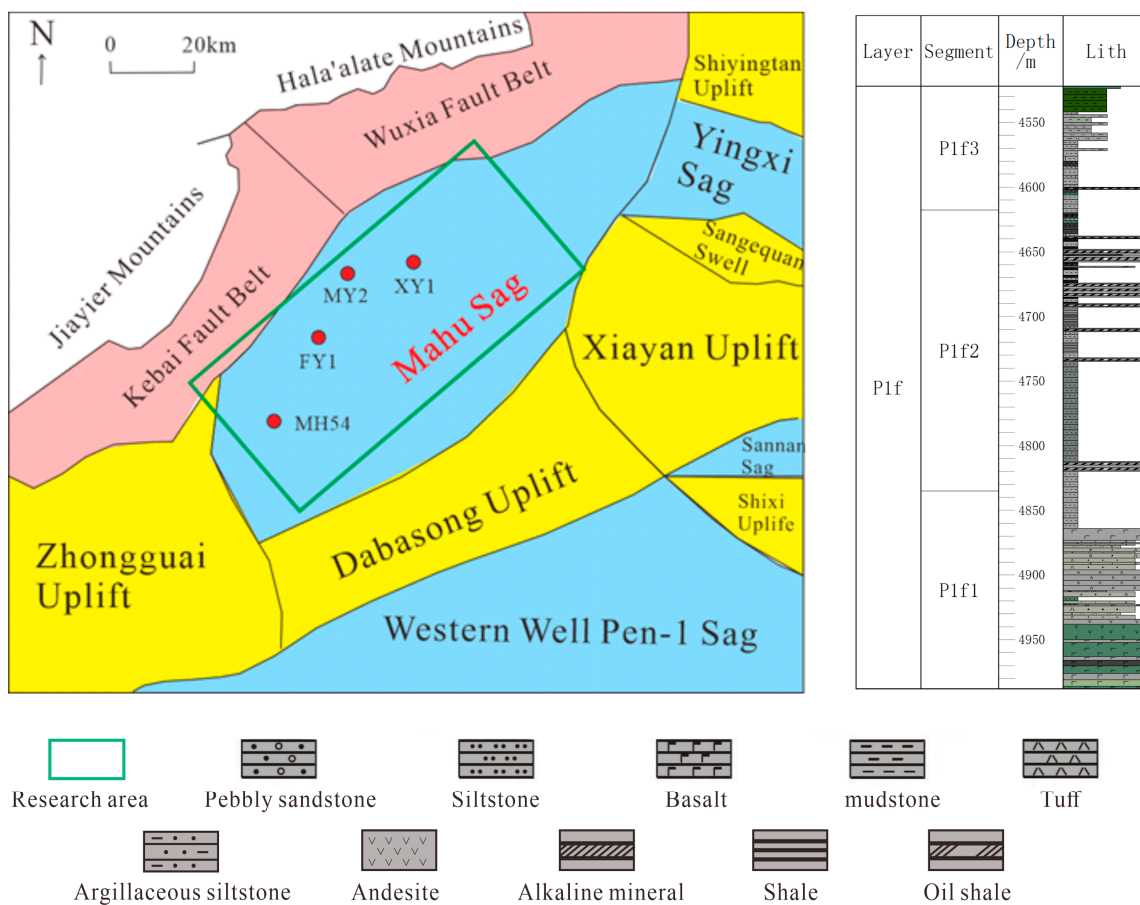


Figure 1. Structural position and stratigraphic column of the Mahu Sag.

3. Main Mineral Types and Alkaline Mineral Characteristics in the Research Area

The lithology of the Fengcheng Formation in the Mahu Sag is complex, with a diverse range of mineral types and significant variations in their content. The mineral characteristics of this formation bear similarities to the Green River Formation of the Eocene Uinta Basin in the eastern part of Utah, USA [15]. This article selects 29 cores with different alkaline mineral contents for X-ray diffraction experiments to analyze their main mineral composition (Figure 2). It can be seen from the figure that the main minerals of the Fengcheng Formation are feldspar, alkaline minerals, and quartz. Among them, feldspar content is the greatest, ranging from 10% to 70%, and most of them range from 20% to 50%. This coring mainly comes from the Feng-2(P_{1f}^2) segment, so the content of alkaline minerals is relatively high, with a total number of alkaline minerals ranging from 0 to 60%. Among the alkaline minerals, the content of trona is the highest, and some cores have a trona content of 100%, which is pure trona cores; the quartz content mainly ranges from 5% to 30%; simultaneously developing a small amount of pyrite, its content is generally less than 10%.

The main alkaline minerals developed in the Fengcheng Formation of the Mahu Sag include trona, sodium bicarbonate (soda ash), eitelite, northupite, shortite, and reedmergnerite, with a small amount of gypsum and anhydrite locally developed [16,17]. This article mainly divides alkaline minerals into four categories based on their chemical composition: trona minerals, eitelite minerals, shortite minerals, and silicate borate minerals. Trona minerals mainly consist of trona (Na_2CO_3) and soda ash ($Na_5H_3(CO_3)_4$); eitelite minerals mainly include eitelite ($MgNa_2(CO_3)_2$) and northupite ($Na_3Mg(CO_3)_2Cl$); shortite minerals mainly include shortite ($Ca_2Na_2(CO_3)_3$), gaylussite ($CaNa_2(CO_3)_2 \cdot 5H_2O$), and calcium water alkali ($CaNa_2(CO_3)_2 \cdot 2H_2O$); and the main mineral of silicate borate is reedmergnerite ($Na(BSi_3O_8)$).

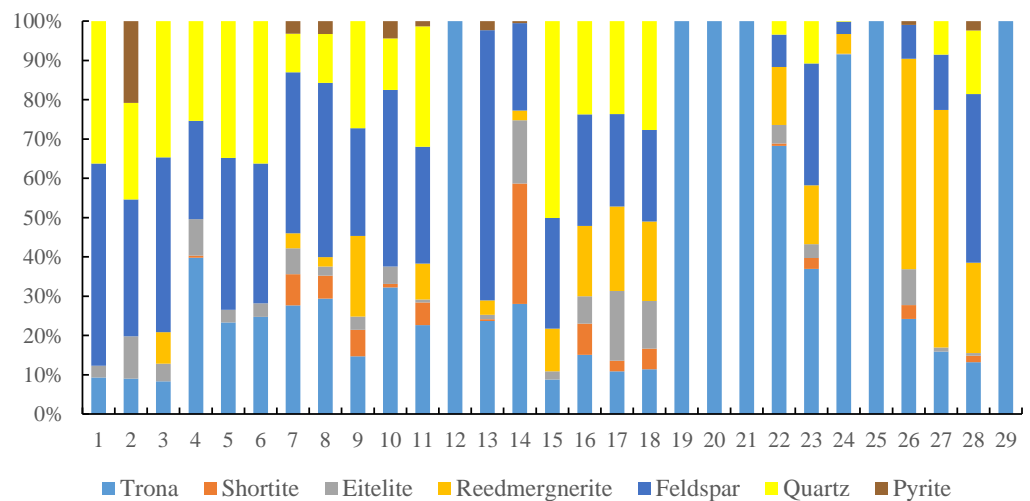


Figure 2. Mineral types and content of Fengcheng Formation in Mahu Sag.

Trona minerals have low hardness and are prone to deliquescence in the air. They are highly soluble in water and react with fluids in the formation to form corrosion pores (Figure 3a); eitelite minerals are slightly soluble in water and easily soluble in acids, appearing colorless, and transparent in thin sections, often coexisting with shortite minerals [18] (Figure 3b); shortite minerals are insoluble in water and slowly dissolve in acid, exhibiting polychromism under the microscope and having a glassy luster (Figure 3c); and the main mineral of silicate borate is reedmergnerite, which is insoluble in both water and acid and has a hardness greater than that of a knife (Figure 3d).

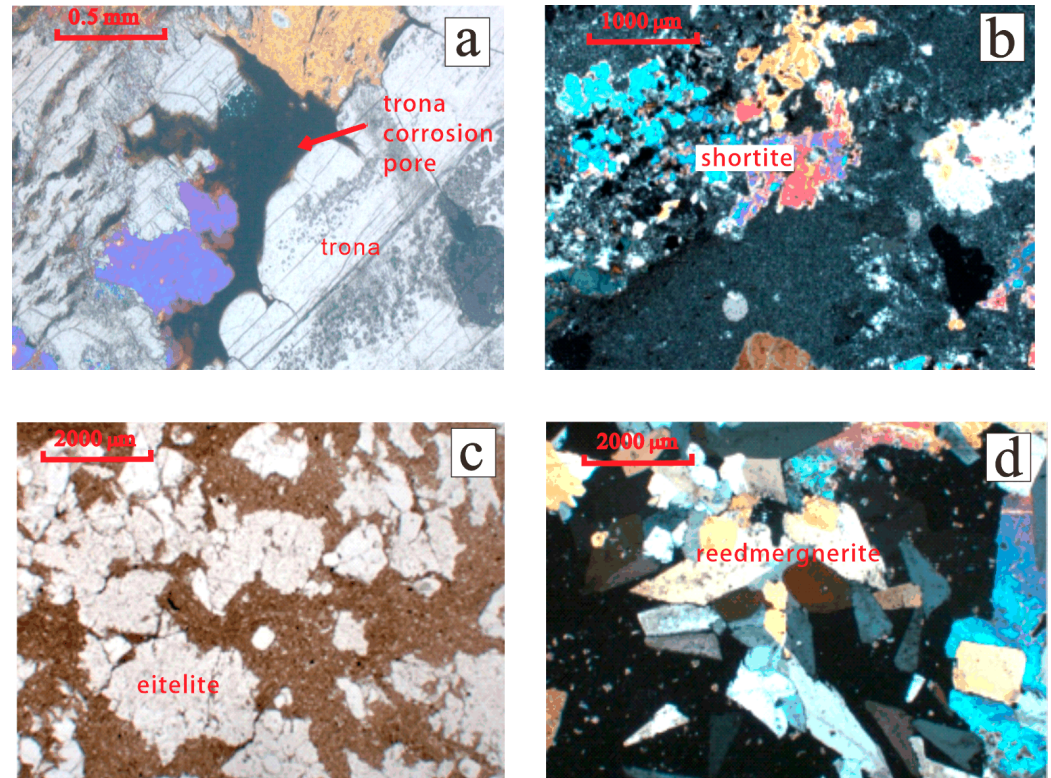


Figure 3. Microscopic characteristics of alkaline minerals from Fengcheng Formation in Mahu Sag. (a) Trona dissolution, Feng 26 well, 3304.1m. (b) Microscopic characteristics of shortite, Maye 2 well, 4771.4 m. (c) Microscopic characteristics of eitelite, Fengnan 5 well, 3534.5 m. (d) Microscopic characteristics of reedmergnerite, Aike 1 well, 5477.6 m.

4. Methodology

4.1. Improved Petrophysical Volume Modeling

Strata is composed of different minerals and fluids in pores, with minerals mainly including skeleton minerals and clay minerals, and fluids in pores mainly including oil, gas, and water. The skeleton minerals of conventional reservoirs mainly include minerals such as feldspar, quartz, dolomite, calcite, pyrite, mica, etc. Therefore, the rock volume physical model used in the inversion of conventional reservoir minerals mainly targets the common minerals mentioned above (Figure 4a). The Fengcheng Formation in the Mahu Depression is an alkaline lake sedimentary system, with a large amount of alkaline minerals developed. Previous X-ray diffraction experiments showed that the main minerals in the study area were quartz, feldspar, pyrite, shortite, eitelite, reedmergnerite, and trona. However, the existing rock volume physical models do not include alkaline minerals such as trona, shortite, eitelite, and reedmergnerite. Therefore, improvements were made to the existing rock volume physical model, and based on the results of previous X-ray diffraction experiments, shortite, eitelite, reedmergnerite, and trona were added to the rock volume physical model to establish a rock volume physical model suitable for alkali-bearing shale oil reservoirs (Figure 4b).

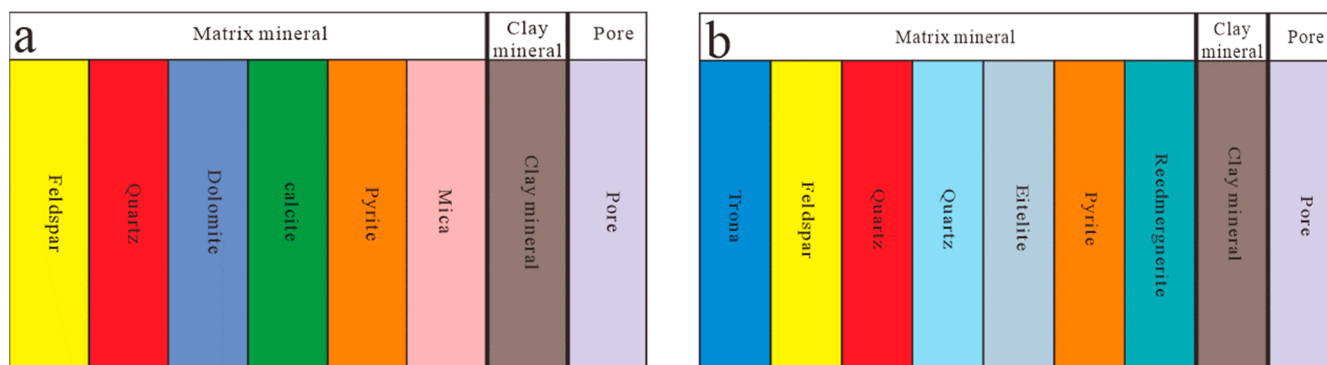


Figure 4. Schematic diagram of rock volume physical model: (a) conventional reservoir rock volume physical model and (b) physical model of rock volume in alkali-bearing shale oil reservoirs.

The element content in the formation is closely related to the mineral content. Previous studies have determined the matrix relationship between the minerals in the formation and the formation elements through a large number of experiments. Herron et al. used the coefficient matrix inversion method as the basis to calculate the mineral content, using the dry weight of elements [19,20]. Based on the physical model of alkali-bearing rock volume mentioned above, combined with the method of coefficient matrix, a mineral content matrix equation is established for alkali-bearing shale oil reservoirs, and constraint conditions are added based on the actual lithology characteristics of the formation (Equation (1)).

$$\begin{pmatrix} a_{11} & a_{12} & a_{13} & \cdots & a_{1m} \\ a_{21} & a_{22} & a_{23} & \cdots & a_{2m} \\ a_{31} & a_{32} & a_{33} & \cdots & a_{3m} \\ \vdots & \vdots & \vdots & \ddots & \vdots \\ a_{n1} & a_{n2} & a_{n3} & \cdots & a_{nm} \end{pmatrix} \times \begin{pmatrix} x_1 \\ x_2 \\ x_3 \\ \vdots \\ x_m \end{pmatrix} = \begin{pmatrix} b_1 \\ b_2 \\ b_3 \\ \vdots \\ b_n \end{pmatrix} \quad (1)$$

where m is the number of skeleton components; x_j is the content of the j component ($j = 1, 2, \dots, m$), decimal; a_{ij} is the response equation coefficient of the i curve of the j mineral component ($i = 1, 2, \dots, n$; n is the number of selected log curves); b_i is the value of the i formation element log curve, decimal or 10^{-6} . Its constraint condition is $\sum_{j=1}^n x_j = c$, where c is a constant; $0 \leq x_j \leq x_{max,j}$, $x_{max,j}$ is the maximum content of the j component found in

the formation. The response coefficients of each mineral in Litho Scanner Log are shown in Table 1.

Table 1. Response coefficients of mineral components.

	Al Weight Fraction (lbf/lbf)	Ca Weight Fraction (lbf/lbf)	Fe Weight Fraction (lbf/lbf)	Mg Weight Fraction (lbf/lbf)	K Weight Fraction (lbf/lbf)	Si Weight Fraction (lbf/lbf)	S Weight Fraction (lbf/lbf)
Quartz	0.0000	0.0000	0.0000	0.0000	0.0000	0.4675	0.0000
Feldspar	0.0990	0.0010	0.0010	0.0010	0.0510	0.3000	0.0000
Pyrite	0.0000	0.0000	0.4655	0.0000	0.0000	0.0000	0.5345
Shortite	0.0000	0.2614	0.0000	0.0000	0.0000	0.0000	0.0000
Eitelite	0.0000	0.0000	0.0000	0.0960	0.0000	0.0000	0.0000
Reedmergnerite	0.0000	0.0000	0.0000	0.0000	0.0000	0.3417	0.0000

When the number of selected element curves is equal to the number of mineral components in the formation ($m = n$), the mineral dry weight can be calculated using the inverse matrix solution method, using Equation (2):

$$\begin{pmatrix} x_1 \\ x_2 \\ x_3 \\ \vdots \\ x_m \end{pmatrix} = \begin{pmatrix} \frac{a_{11}^*}{\det(A)} & \frac{a_{12}^*}{\det(A)} & \frac{a_{13}^*}{\det(A)} & \cdots & \frac{a_{1n}^*}{\det(A)} \\ \frac{a_{21}^*}{\det(A)} & \frac{a_{22}^*}{\det(A)} & \frac{a_{23}^*}{\det(A)} & \cdots & \frac{a_{2n}^*}{\det(A)} \\ \frac{a_{31}^*}{\det(A)} & \frac{a_{32}^*}{\det(A)} & \frac{a_{33}^*}{\det(A)} & \cdots & \frac{a_{3n}^*}{\det(A)} \\ \vdots & \vdots & \vdots & \ddots & \vdots \\ \frac{a_{m1}^*}{\det(A)} & \frac{a_{m2}^*}{\det(A)} & \frac{a_{m3}^*}{\det(A)} & \cdots & \frac{a_{mn}^*}{\det(A)} \end{pmatrix} \times \begin{pmatrix} b_1 \\ b_2 \\ b_3 \\ \vdots \\ b_n \end{pmatrix} \quad (2)$$

where a_{ji}^* is the adjoint matrix coefficient of a_{ij} , and $\det(A)$ is the determinant of the coefficient matrix in Equation (1).

Therefore, the dry weight content of minerals is as follows:

$$h_i = \sum_{j=1}^n \frac{a_{ij}^*}{\det(A)} \times b_j / \sum_{j=1}^m \sum_{i=1}^n \frac{a_{ij}^*}{\det(A)} \times b_i \times 100\% \quad (3)$$

where h_i is the mass percentage of the i -th mineral, %.

When the number of selected element curves is not equal to the number of formation mineral components, based on Equation (1), according to the principle of the least squares method, solving Equation (1) can be transformed into a problem of finding the extreme value (Equation (4)).

$$\begin{cases} \min f(x) = \sum_{i=1}^m \left(\sum_{j=1}^n a_{ij} x_j - b_i \right) \\ \sum_{j=1}^n x_j = c \\ 0 \leq x_j \leq x_{max, j} \end{cases} \quad (4)$$

where $\min f(x)$ is the objective function.

When $n \leq m$, the solution space of Equation (1) is convex and has a unique solution [21]. Therefore, Equation (4) has and only has a minimum value, which is the optimal solution for the content of each mineral component.

4.2. Mineral Assemblage Model

The study area features a diverse range of minerals. When employing a single-mineral model for mineral inversion, the increased number of unknowns amplifies uncertainty, leading to higher errors. Consequently, in practical mineral content inversion, a combined model is often utilized. This involves selecting characteristic elements for different minerals and accurately calculating mineral content based on the principles of optimization. At present, the Elan mineral calculation module from Schlumberger Company, Houston, TX, USA is commonly used for mineral inversion [22,23]: using sulfur element to calculate pyrite content; using silicon dioxide, magnesium carbonate, calcium carbonate, and iron

element content to calculate clay mineral content; using calcium and magnesium element content to determine calcite and dolomite content; and calculating quartz, feldspar, and mica content using other element content [24]. The Fengcheng Formation shale oil reservoir develops a large amount of alkaline minerals, and the Elan module does not include the calculation of alkaline minerals. Therefore, based on the Elan module and combined with the characteristic elements in alkaline minerals, as is shown in Figure 5, this article adopts a combination model method to establish 4 mineral interpretation models: Shorttite Model (feldspar, quartz, pyrite, and shorttite), Eitelite Model (feldspar, quartz, pyrite, and eitelite), Reedmergnerite Model (feldspar, quartz, pyrite, and reedmergnerite), and Trona Model. By combining and calculating 4 models, mineral inversion of alkali-bearing shale oil reservoirs is achieved.

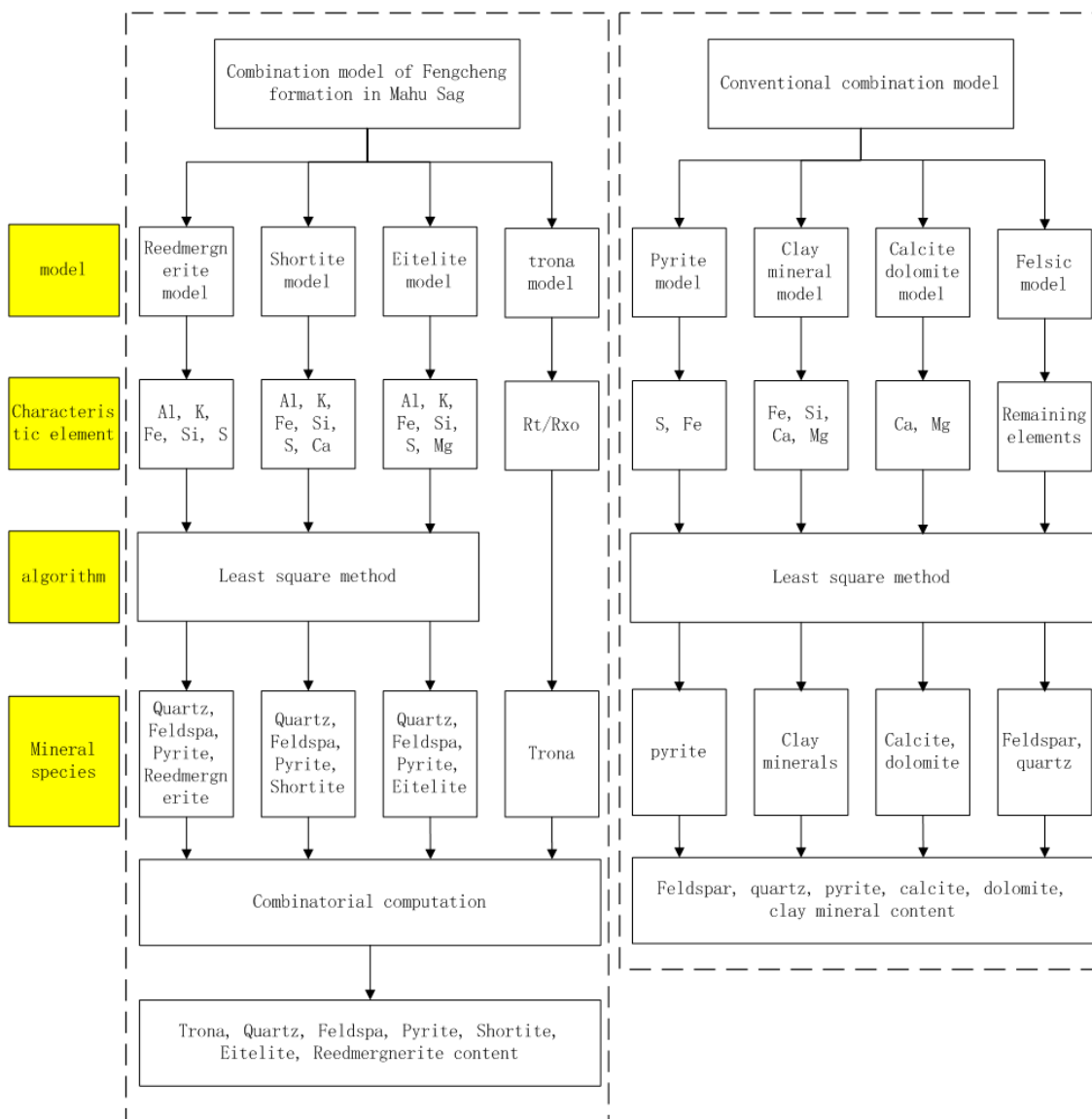


Figure 5. Comparison of calculation flowcharts of mineral association models.

The main mineral in alkaline minerals is trona, which is easily soluble in mud during actual production, resulting in wellbore collapse. However, the radial detection depth of Litho Scanner Log is only 23 cm [25], and the Litho Scanner Log data in trona development intervals are distorted. Trona dissolves in the mud, increasing the mineralization degree of the mud, leading to a decrease in the mud resistivity. Therefore, in layers with high trona content, the ratio of the original formation resistivity to the flushing resistivity is relatively

large. As shown in Figure 6, the red dot in the figure is the content of trona in the core, and the black trendline is the relationship between the ratio of virgin zone resistivity to flush zone resistivity and trona content established by sigmoid function [26] (Equation (5)).

$$V = \begin{cases} \frac{421}{1.97 + 2.36 \times (R_t/R_{x0})^{-0.43}} - 100 & R_t/R_{x0} \leq 752.21 \\ 100 & R_t/R_{x0} > 752.21 \end{cases} \quad (5)$$

where V is the trona content, %; R_t is the original formation resistivity, $\Omega \cdot m$; and R_{x0} is the formation resistivity of the flushing zone, $\Omega \cdot m$.

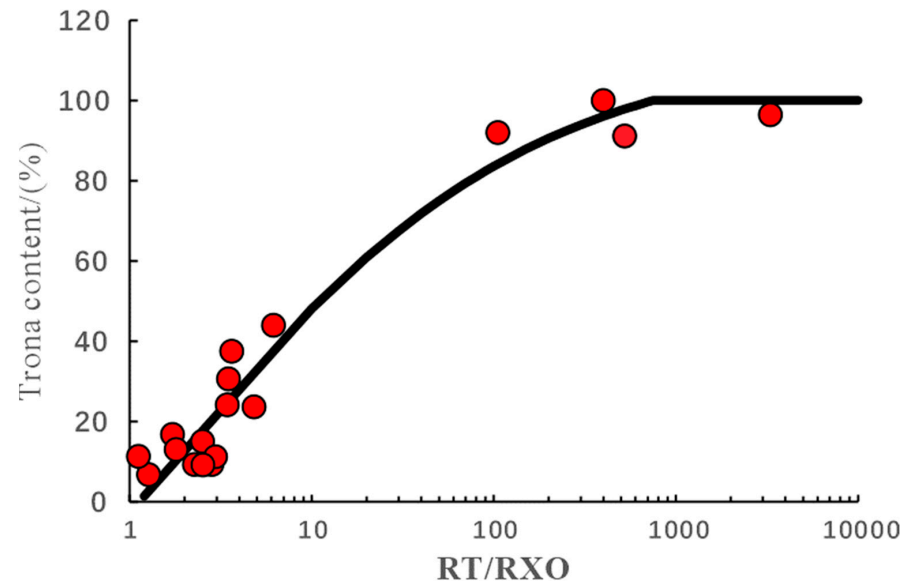


Figure 6. Prediction model for trona content.

In addition to the prediction model of trona content, based on the characteristic elements of minerals, we established the Shortite Model (feldspar, quartz, pyrite, and shortite), Eitelite Model (feldspar, quartz, pyrite, and eitelite), and Reedmergnerite Model (feldspar, quartz, pyrite, and reedmergnerite). We selected different elements to perform mineral inversion on different models, and the dry weight of aluminum, potassium, iron, silicon, sulfur, and calcium elements was inverted using the Shortite Model mineral. The dry weight of aluminum, potassium, iron, silicon, sulfur, and magnesium elements was inverted using the Eitelite Model mineral. The dry weight of aluminum, potassium, iron, silicon, and sulfur elements was used to invert the model minerals of Reedmergnerite Model. Feldspar, quartz, feldspar, pyrite, shortite, eitelite and reedmergnerite were inverted based on the model mineral composition, and the mineral content was ultimately determined by combining the model method. We calculated the element response coefficients of each mineral through its chemical composition (Table 1) and performed mineral inversion for each model, using the least squares method.

We calculated and combined the mineral inversion results of the four models (Equation (6)) and obtained the final mineral content.

$$\begin{aligned} W_t &= M_t/100 \\ W_e &= (100 - M_t) \times M_e/100 \\ W_s &= (100 - M_t) \times M_s/100 \\ W_r &= (100 - M_t) \times (100 - M_e - M_s) \times M_r/100 \\ W_q &= (100 - M_t) \times (100 - M_e - M_s) \times M_q/100 \\ W_f &= (100 - M_t) \times (100 - M_e - M_s) \times M_f/100 \\ W_p &= (100 - M_t) \times (100 - M_e - M_s) \times M_p/100 \end{aligned} \quad (6)$$

where W_t is the content of trona, %; W_r is the content of reedmergnerite, %; W_e is the content of eitelite, %; W_s is the content of shortite, %; W_q is the content of quartz, %; W_f is the content of feldspar, %; W_p is the content of pyrite, %; M_t is the trona content calculated by the Trona Model, %; M_r is the content of reedmergnerite calculated by the Reedmergnerite Model, %; M_e is the content of eitelite calculated by the Eitelite Model, %; M_s is the content of shortite calculated by the Shortite Model, %; M_q is the quartz content calculated by the Reedmergnerite Model, %; M_f is the feldspar content calculated by the Reedmergnerite Model, %; and M_p is the pyrite content calculated by the Reedmergnerite Model, %.

4.3. Oxide Closure Model

Litho Scanner Log can measure the energy spectrum of elements in formation. During the actual log process, the element energy spectrum measured by the instrument is usually regarded as a linear combination of the standard spectrum of different elements [27,28]. The element energy spectrum count rate can be expressed according to Equation (7):

$$c_i = \sum_{j=1}^m a_{ij}y_j + \varepsilon_i \quad (7)$$

where c_i is the energy of the i spectrum of the element energy spectrum measured by the Litho Scanner tool, MeV; a_{ij} is the standard energy magnitude of the j -th element in the i -th spectrum, MeV; y_j is the yield of the j -th element, %; ε_i is the error, MeV; m is the total number of elements, dimensionless; and n is the total number of spectra, dimensionless.

The yields of different elements can be determined through the weighted least square method [29,30]. As indicated in Equation (8), the yields for each element can be calculated by minimizing R.

$$R = \sum_{i=1}^n w_i \varepsilon_i^2 = \sum_{i=1}^n w_i (c_i - \sum_{j=1}^m a_{ij}y_j) \quad (8)$$

The yield of Al, Si, Ca, Fe, K, Na, Mg, S, Ti, and Gd was used to calculate the dry weight of formation elements based on the oxide closure model. The core idea of the oxide closure model is that all minerals in the formation are composed of oxide and carbonate, and the sum of percentage of oxide and carbonate content is 1; the sum of the mass percentage of all elements is also 1 [31]. As shown in Equations (9) and (10), the content of major elements in the rock skeleton can be calculated by the yield of each element and the oxide index. The oxide closure model is used to normalize the oxide and carbonate minerals, which overcomes the problem that it is difficult for instruments to measure the yield of carbon and oxygen [32]. Figures 7 and 8 are the results of element dry weight calculation of Litho Scanner Log in the Fengyun 1 well and Mahu 54 well, respectively.

$$W_i = F(y_j/s_j) \quad (9)$$

$$1 = F\left(\sum_j \frac{x_j \times y_j}{s_j}\right) \quad (10)$$

where W_i is the mass percentage of element i , %; F is the closure factor to be determined for each depth point; x_j is the oxide index of formation element j ; y_j is the relative yield of the measured formation element j , %; and s_j is the relative detection sensitivity factor of formation element j , $g \times s$.

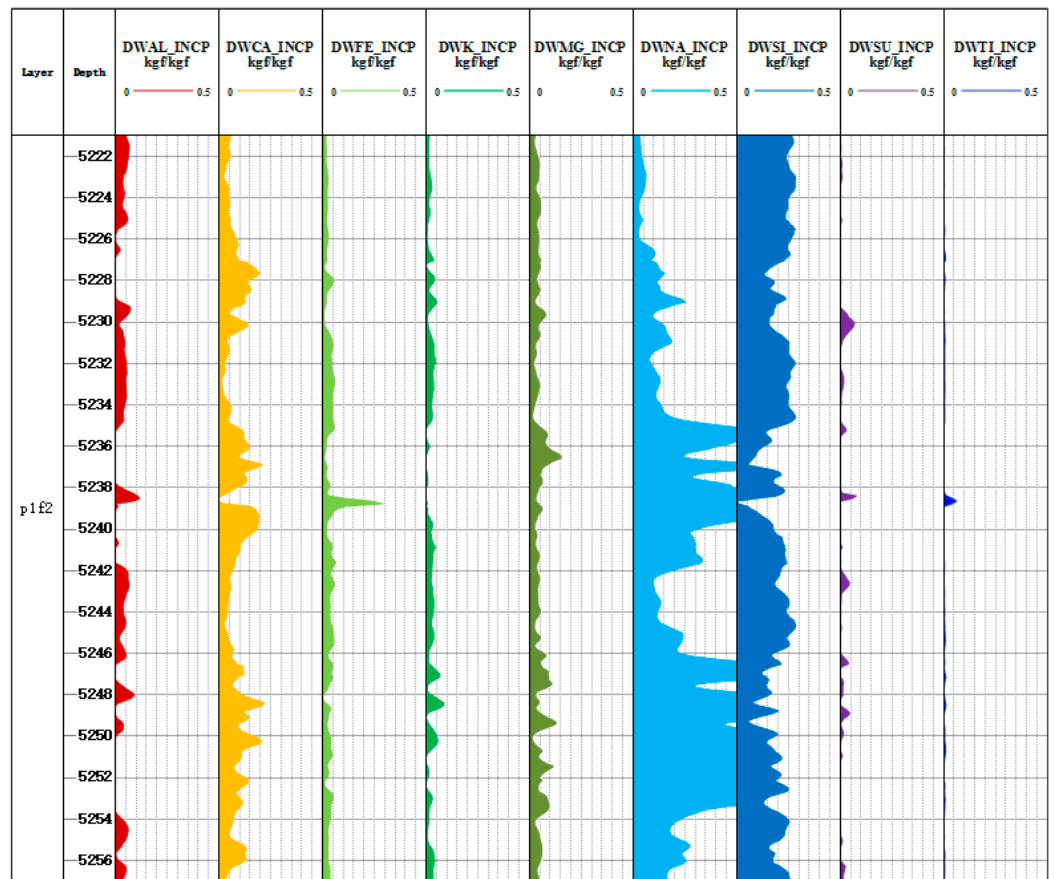


Figure 7. Dry weight diagram of elements in Well Fengyun 1.

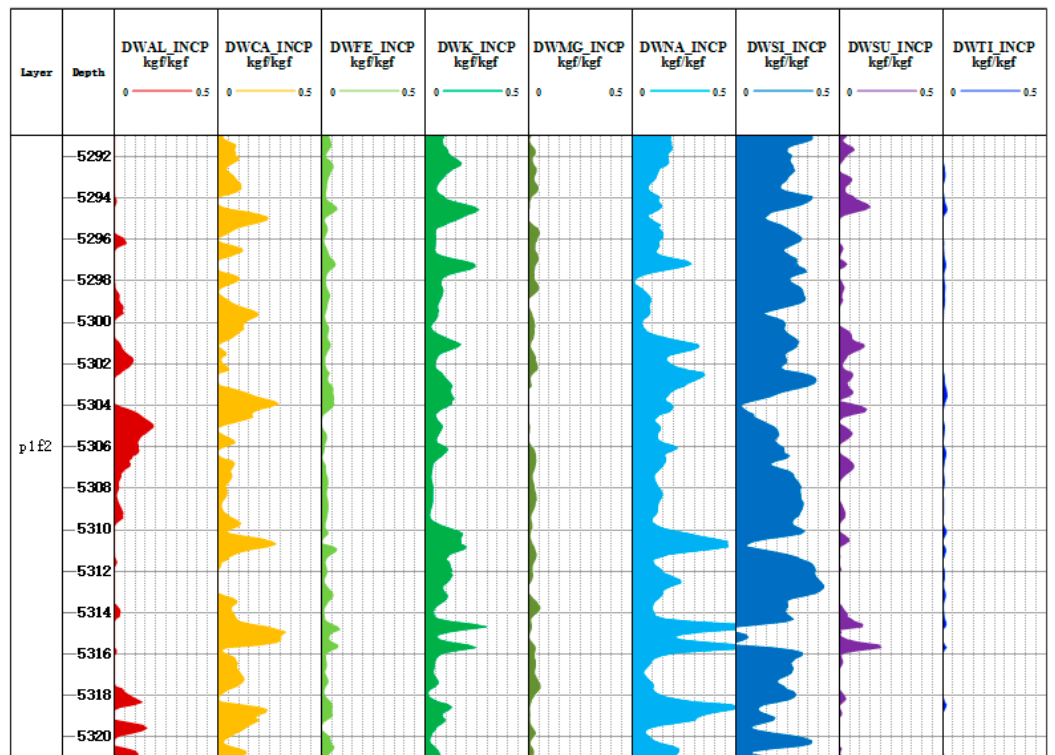


Figure 8. Dry weight diagram of elements in Well Mahu 54.

5. Application and Evaluation of Mineral Inversion

5.1. Litho Scanner Log Mineral Inversion Results

If we compare the mineral inversion results with the core X-ray diffraction experimental results, as shown in Figures 9 and 10, we can see that the third trace is the inverted mineral profile. From the beginning of the 1st to 10th tracks, they are composed of a shortite component, eitelite component, reedmergnerite component, feldspar component, quartz component, pyrite component, and trona component. This is a gradient color fill, the darker the color the more mineral content. Pink fill channel is shortite; reddish brown fill channel is eitelite; brown fill channel is reedmergnerite; purple fill channel is feldspar; yellow fill channel is quartz; dark green fill channel is pyrite; blue fill channel is trona. The correlation coefficient between the mineral inversion results of Litho Scanner Log and the results of core X-ray diffraction experiment is 0.847. Most of the inversion results have an absolute error of no more than 20% from the core data, with the highest amount of trona among the core minerals with an absolute error of more than 20% (Figure 11). The mineral inversion results have good consistency with the core experiment results.

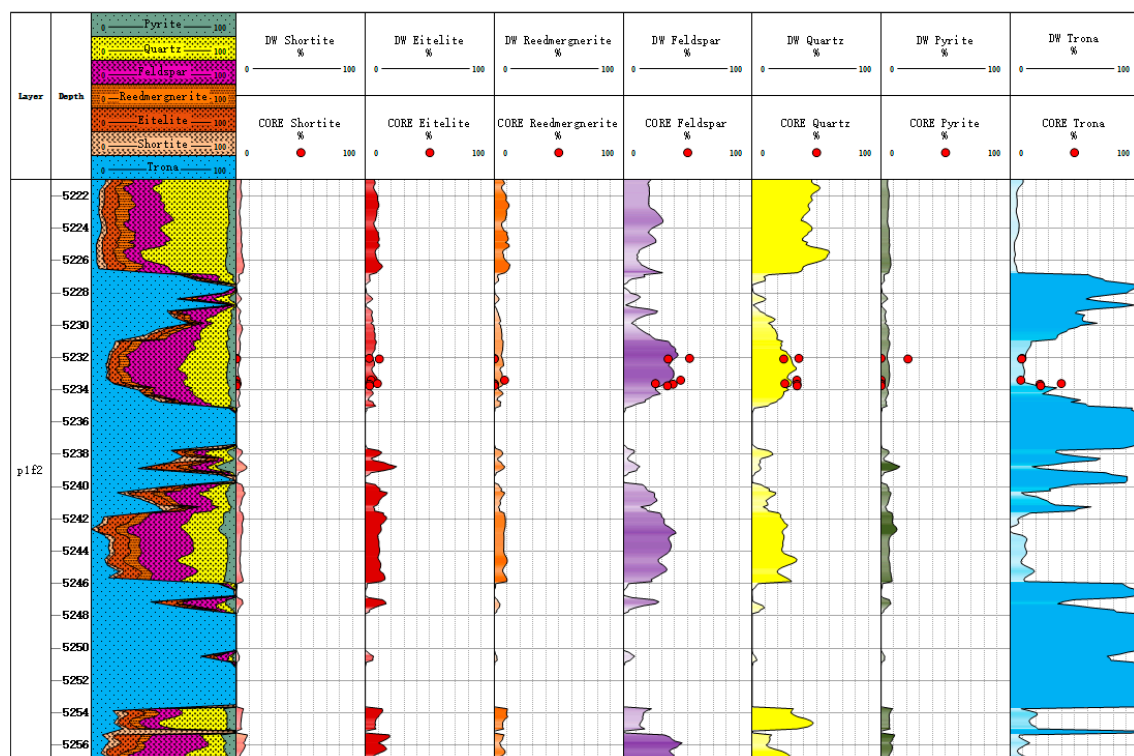


Figure 9. Mineral content map of Well Fengyun 1.

5.2. Conventional Log Mineral Inversion Results

Litho Scanner Log is expensive, and its widespread application will increase exploration costs. Therefore, the mineral inversion results of Litho Scanner Log are calibrated with conventional log data to achieve mineral inversion of alkali-bearing shale oil reservoirs based on conventional log data. We improved the rock volume physical model of conventional log by adding shortite, eitelite, reedmergnerite, and trona to the rock volume physical model and established a rock volume physical model of alkali-bearing ore based on conventional log data.

We then selected density curve, neutron porosity curve, acoustic time difference curve, natural gamma curve, deep resistivity curve, and flushing resistivity curve for mineral inversion of alkali-bearing shale oil reservoirs. By using the principle of the least squares method and combining the above combination model for mineral inversion, the mineral

inversion of alkali-bearing shale oil reservoirs using conventional log data was achieved. The selected mineral log parameters in this article are shown in Table 2.

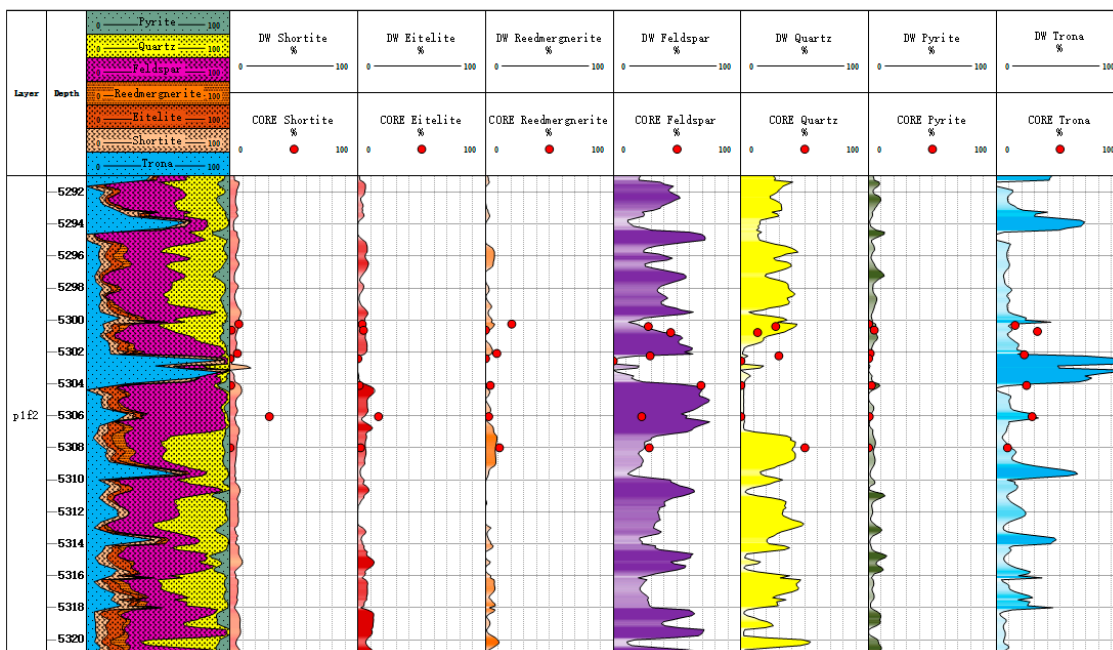


Figure 10. Mineral content map of Well Mahu 54.

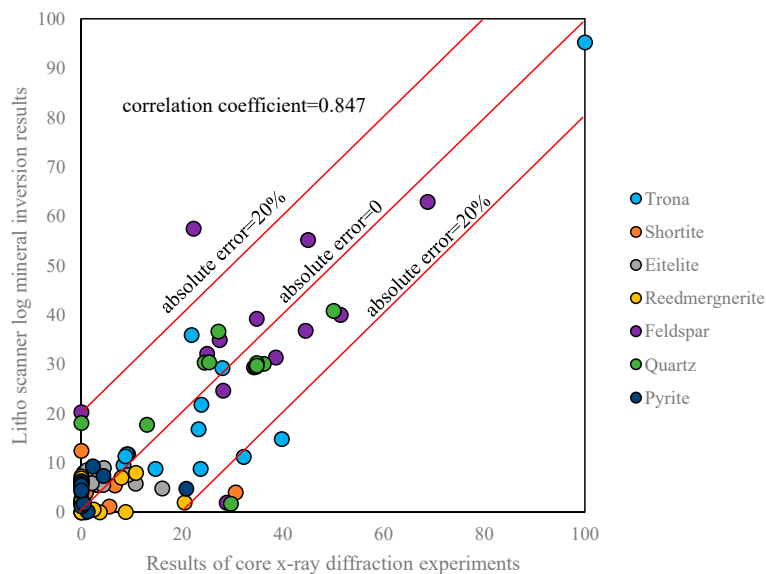


Figure 11. Comparison of mineral inversion and core X-ray diffraction experiments.

Table 2. Response coefficients of the conventional mineral log.

	Bulk Density (g/cm ³)	Neutron Porosity (v/v)	Compressional Slowness (us/ft)	UI GR (gAPI)
Quartz	2.65	−0.03	56	30
Feldspar	2.57	0.02	60	170
Pyrite	4.99	−0.03	39	0
Shortite	2.63	0.11	53	2
Eitelite	2.46	0.30	53	30
Reedmergnerite	2.74	0.18	60	5

Except for trona and feldspar, the average absolute error of other minerals is less than 10%, which has a high coincidence rate and meets the needs of fine log interpretation.

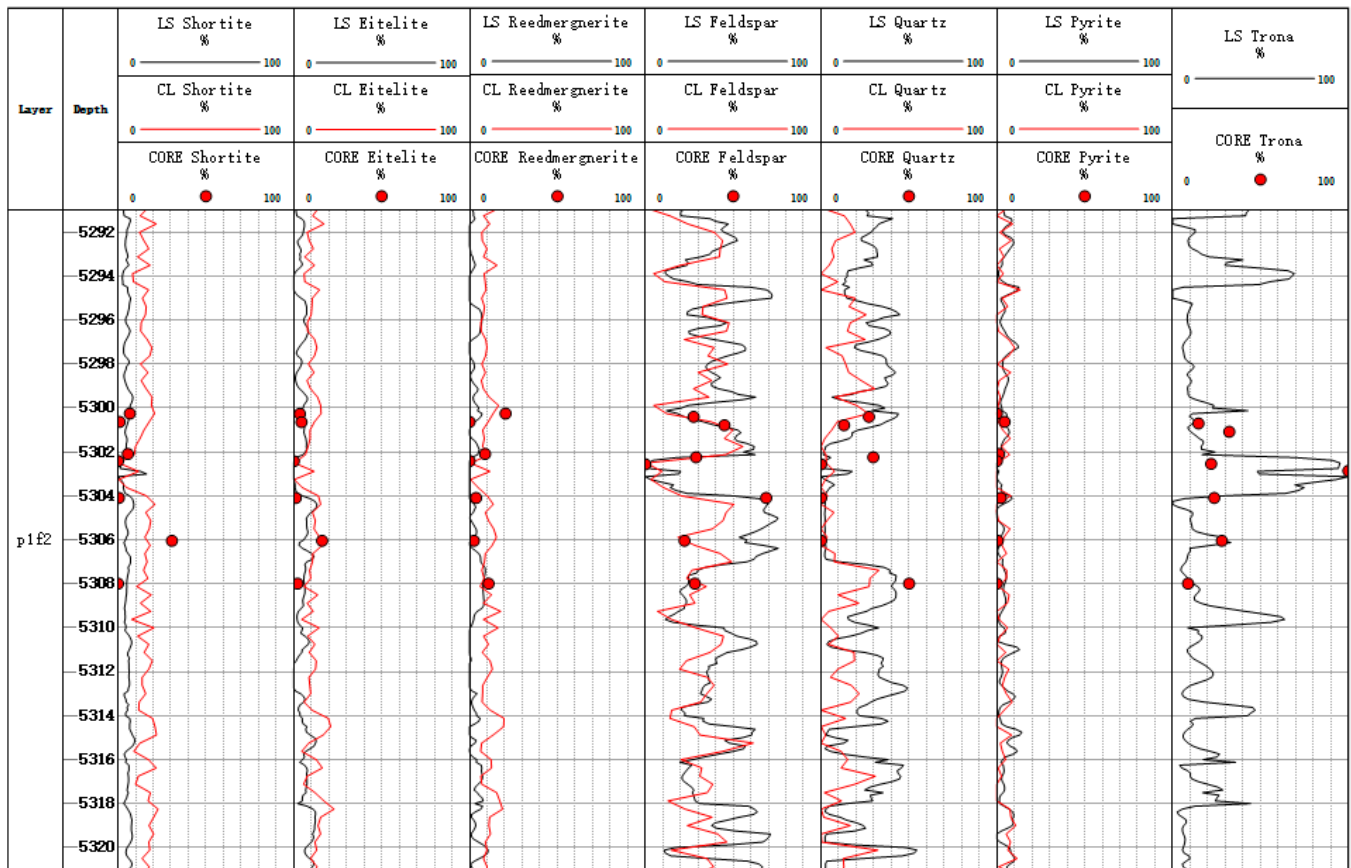


Figure 13. Mineral inversion map of conventional log in Well Mahu 54.

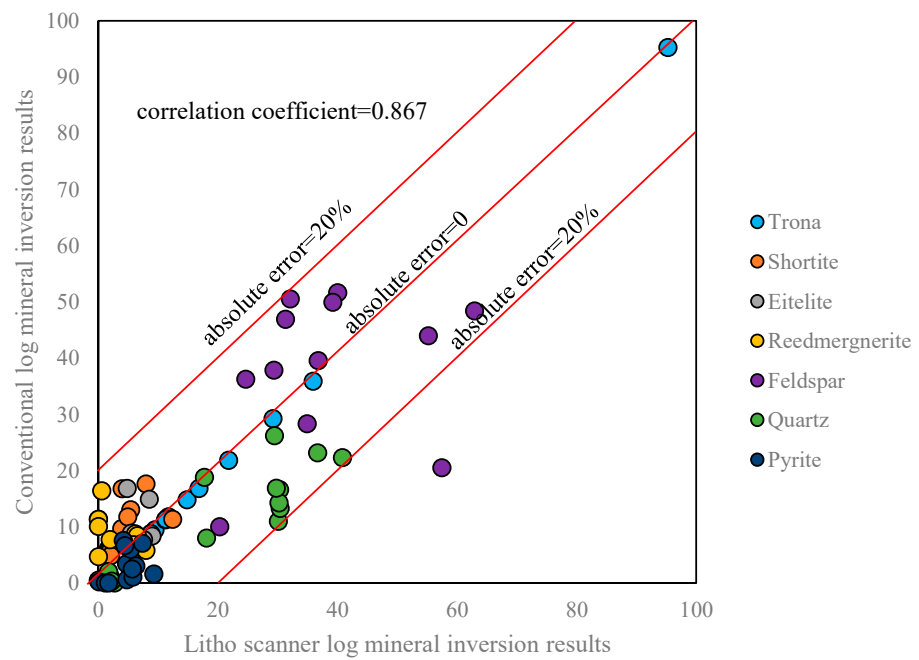


Figure 14. Comparison of mineral inversion results from conventional log and Litho Scanner Log.

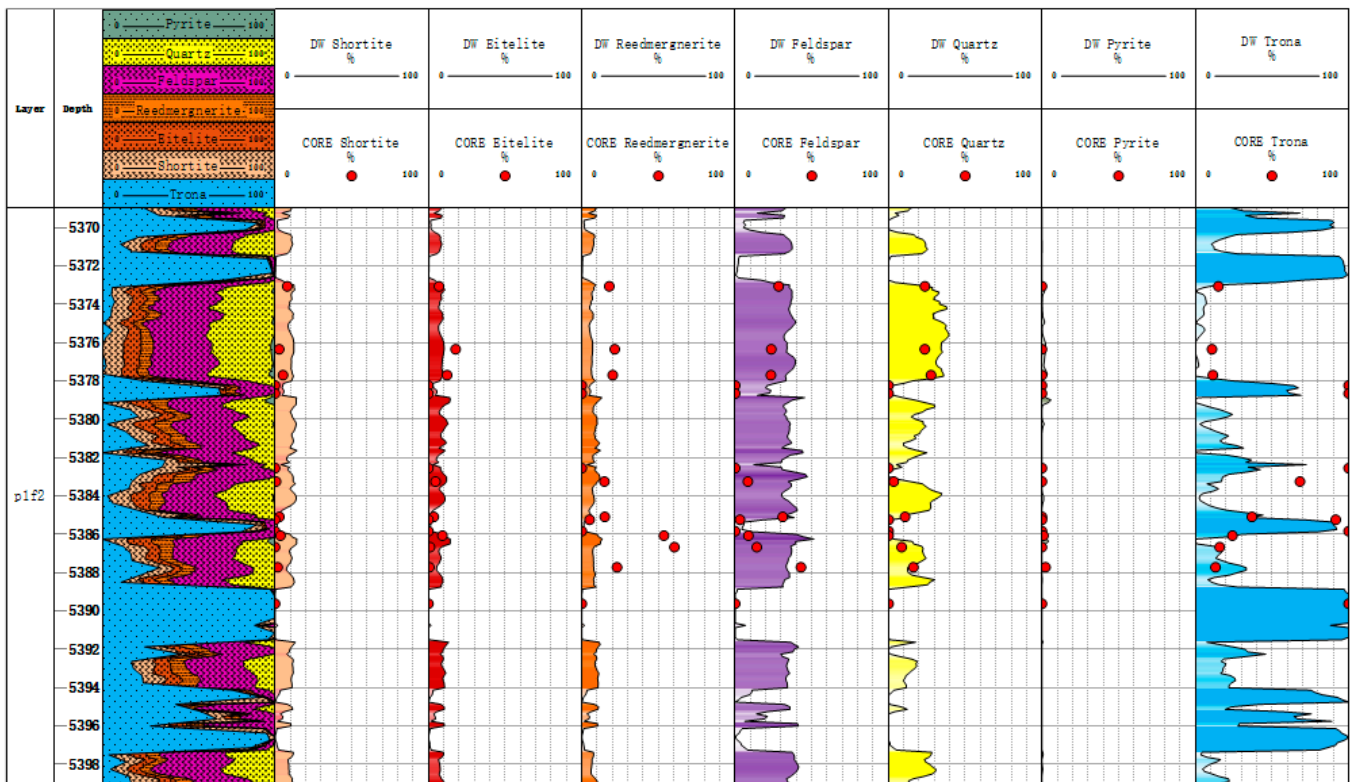


Figure 15. Comparison between mineral content inversion results and XRD experimental results of Well Xiayun 1.

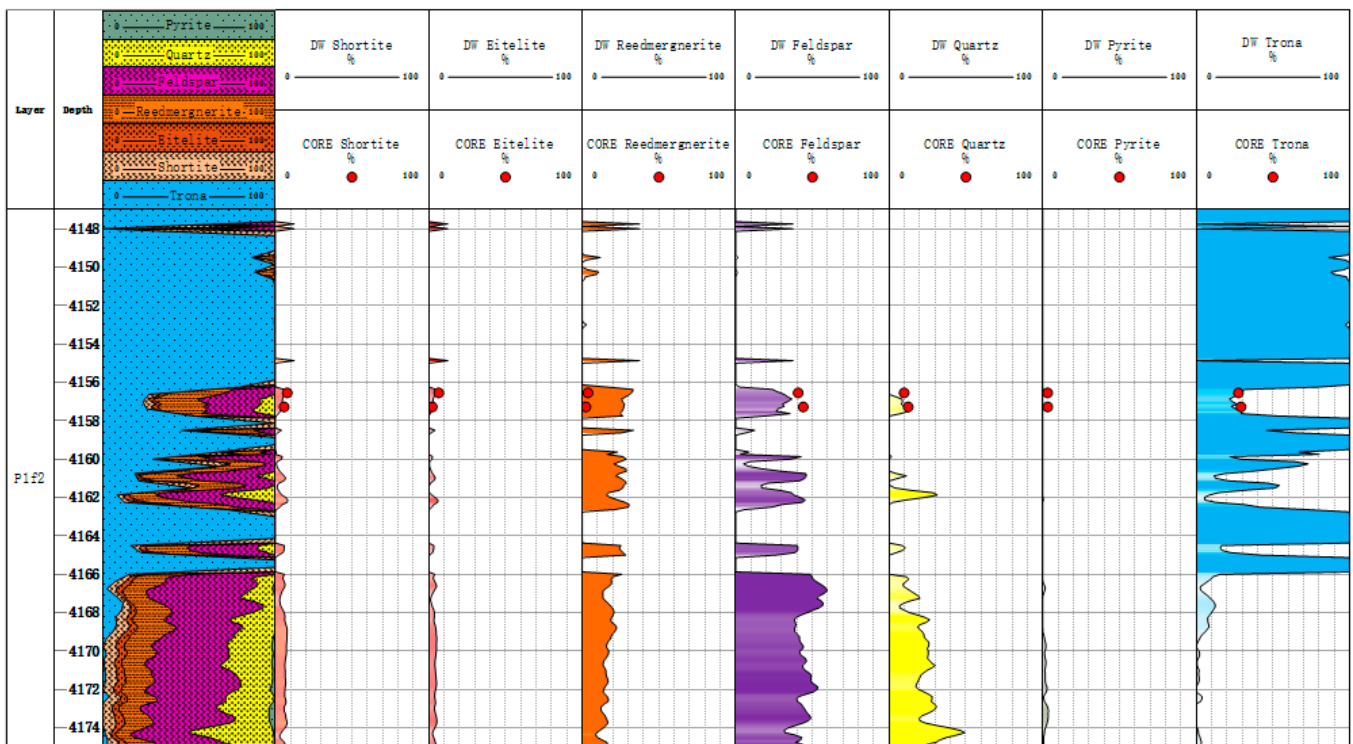


Figure 16. Comparison of mineral content inversion results and XRD experimental results in Well Maye 2.

Table 3. Absolute error analysis table.

Sample Number	Trona		Shortite		Eitelite		Reedmergnerite		Feldspar		Quartz		Pyrite	
	Inversion Content/%	Core Content/%	Inversion Content/%	Core Content/%	Inversion Content/%	Core Content/%	Inversion Content/%	Core Content/%	Inversion Content/%	Core Content/%	Inversion Content/%	Core Content/%	Inversion Content/%	Core Content/%
1	7.7	15.1	12.1	8.0	9.3	6.9	7.6	17.9	34.6	28.4	28.9	23.7	0.0	0.0
2	0.0	10.9	12.2	2.7	10.4	17.7	7.1	21.5	33.9	23.5	34.0	23.6	2.5	0.0
3	0.0	11.4	12.2	5.2	10.4	12.2	6.5	20.2	32.6	23.3	38.8	27.7	3.7	0.0
4	64.1	100.0	5.1	0.0	2.8	0.0	5.1	0.0	20.8	0.0	2.1	0.0	0.0	0.0
5	63.2	100.0	5.3	0.0	5.2	0.0	5.3	0.0	20.2	0.0	0.8	0.0	0.0	0.0
6	37.7	100.0	8.9	0.0	7.9	0.0	8.9	0.0	30.4	0.0	6.1	0.0	0.0	0.0
7	17.1	68.3	11.9	0.6	11.7	4.6	10.3	14.8	34.6	8.3	14.2	3.4	0.3	0.0
8	25.8	36.9	4.4	2.8	1.8	3.5	8.6	15.0	44.1	31.0	15.3	10.8	0.0	0.0
9	71.0	91.6	3.5	0.0	1.9	0.1	4.1	5.1	18.6	3.1	0.7	0.1	0.0	0.0
10	94.6	100.0	0.8	0.0	0.8	0.0	0.8	0.0	3.1	0.0	0.0	0.0	0.0	0.0
11	32.7	24.2	9.6	3.6	9.6	9.1	10.0	53.6	41.3	8.6	0.0	0.0	0.0	1.0
12	13.4	15.9	11.4	0.0	8.6	1.0	8.1	60.4	36.4	14.1	22.0	8.5	0.1	0.0
13	33.2	13.2	7.7	1.7	5.5	0.6	7.8	23.0	33.2	42.9	12.5	16.2	0.0	2.4
14	100.0	100.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0	0.0
15	27.7	27.5	6.2	7.8	4.7	6.6	32.1	3.8	29.1	40.7	0.0	9.7	0.2	3.2
16	23.5	29.2	5.1	5.8	3.2	2.3	27.1	2.4	30.7	44.0	10.3	12.4	0.1	3.3
Average	38.2	52.8	7.3	2.4	5.9	4.0	9.3	14.8	27.7	16.8	11.6	8.5	0.4	0.6
Average absolute error	14.5		4.9		1.8		5.5		11.0		3.1		0.2	

6. Discussion

6.1. Comparison of Research Methods

At present, the research on alkaline minerals mainly focuses on qualitative research, while quantitative research has not made great progress. The research methods are mainly thin slice analysis, scanning electron microscopy, X-ray diffraction experiment, and electron probe method. Thin slice analysis and scanning electron microscopy can observe the morphology and diagenesis of alkaline minerals such as dissolution and cementation. Although these methods can accurately identify alkaline minerals qualitatively, they cannot quantitatively evaluate alkaline minerals. An X-ray diffraction experiment and the electron probe method can quantitatively calculate the content of alkaline minerals through a spectral analysis and the content of elements in the sample, but these methods are only limited to the analysis of experimental samples and cannot be applied to the quantitative evaluation of alkaline minerals in the whole well interval.

In the log evaluation of alkaline minerals, formations with high alkaline mineral content exhibit higher virgin zone resistivity and lower flush zone resistivity characteristics. Previous studies have quantitatively calculated the content of trona by utilizing the ratio of virgin zone resistivity to flush zone resistivity. However, this method is only applicable to the quantitative assessment of trona. Therefore, this study, in conjunction with the quantitative evaluation method for trona, utilized Litho Scanner Log and conventional well log data as a basis to quantitatively assess other alkaline minerals (such as shortite, eitelite, and reedmergnerite). Leveraging the characteristic elements of alkaline minerals and employing the least squares method, a combination model was established (Reedmergnerite Model, Eitelite Model, Shortite Model, and Torna Model). Through the combined computation of these models, a quantitative assessment of alkaline minerals was achieved.

6.2. Research Significance

Some wells in the Fengcheng Formation of Mahu Sag, Junggar Basin, China, have been drilled into the center of alkaline lakes, where alkaline minerals are highly developed. These minerals, being easily soluble, pose a risk of blockage in actual production engineering, significantly impeding development progress. During the oil testing process in the Fengcheng Formation, industrial oil flow was achieved. The oil-producing interval primarily comprises the shale oil interval between the layers of alkaline minerals. Alkaline minerals themselves represent an important industrial resource with substantial development value. Hence, there is a need for a quantitative evaluation of the alkaline

mineral content. The research outcomes of this paper will contribute to breakthroughs in oil and gas exploration in the Fengcheng Formation of the Mahu Depression and establish a foundation for the enhancement of production extraction technology in alkaline mineral reservoirs in subsequent phases.

7. Conclusions

(1) Fengcheng Formation in Mahu Sag, China, is an alkali-lacustrine facies deposit, which develops a large number of sodium-rich alkaline minerals, which have a significant impact on the accurate evaluation of mineral content in this area. Combined with the results of core X-ray diffraction experiments, the alkaline minerals of the Fengcheng Formation in the Mahu Depression were classified into four groups according to their chemical element compositions: trona minerals (trona and soda ash), eitelite minerals (eitelite and northupite), shortite minerals (shortite, gaylussite, and calcium water alkali), and silicate borate minerals (reedmergnerite). The trona content is the highest among alkaline minerals.

(2) Previous studies on alkaline minerals were mainly qualitative, and quantitative studies on alkaline minerals have not yet made much progress. For the alkaline-bearing shale oil reservoir in the Fengcheng Formation, Mahu Sag, improvements have been made to the existing rock volume physical model. Trona, shortite, eitelite, and reedmergnerite have been incorporated into the rock volume physical model to establish a model suitable for the alkaline-bearing shale oil reservoir. Leveraging Litho Scanner Log and considering the characteristic elements of minerals, models for shortite (feldspar, quartz, pyrite, and shortite), eitelite (feldspar, quartz, pyrite, and eitelite), and reedmergnerite (feldspar, quartz, pyrite, and reedmergnerite) were established. The ratio of virgin zone resistivity to flush zone resistivity was utilized to calculate the content of trona. Finally, the mineral content was determined through a combination model method, forming a comprehensive set of mineral inversion methods for shale oil reservoirs containing alkali ores.

(3) Litho Scanner Log is expensive, and its popularity is low, so utilizing Litho Scanner Log mineral inversion results to scale conventional log mineral inversion results to promote region-wide application is difficult. Based on the improved petrophysical volume model, the density curve, neutron porosity curve, acoustic slowness log, natural gamma curve, virgin zone resistivity curve, and flush zone resistivity curve were selected to invert the mineral content of alkali-bearing shale oil reservoir. Using the principle of the least square method and combined with the method of the combined model, the inversion of mineral in alkaline shale oil reservoir from conventional log data was realized. The inversion results are in good agreement with those of the core X-ray diffraction test.

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